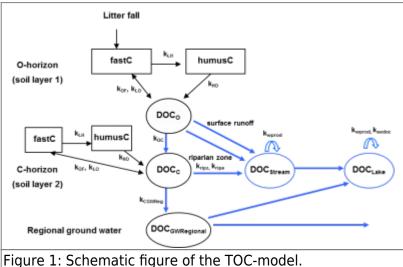
Organic carbon



Source of organic material

Litter fall in the form of plant residues add organic material to HYPE. It increases the levels of fastC in top two layers in soil. The organic carbon addition by litter fall is defined based on crop. Input, resc (kg/ha/yr), gives a daily supplement to the pool during the number of days determined by parameter litterdays.

Links to relevant procedures in the code

Modules (file)	Procedures	
npc soil processes (npc soil proc.f90)	soil carbon processes	

Soil processes

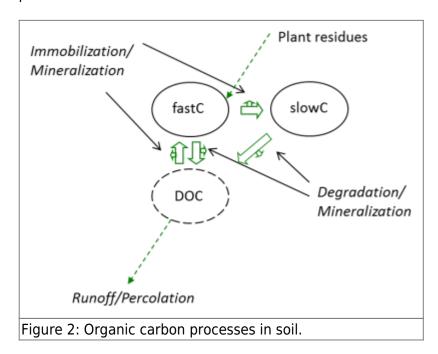
Soil pool initial values

The initial size of organic carbon pools in the soil is dependent on land use and determined by the user. The parameters (humusc1, humusc2, humusc3, fastc1, fastc2, fastc3) give OC content of the three soil layers. The unit for these parameter values is mg/m3. With this information, the pools the size in the different layers are calculated. The model transforms pools into the unit kg/km2 by taking into account the thickness in the layers.

Common functions

Many soil processes depend on temperature and soil moisture. They use the same common functions as nitrogen and phosphorus. Organic carbon soil transformations (production of humusC from fastC, turnover of fastC and turnover of humusC) use the soil moisture function with parameters given by θ_{low} is replaced by the the user instead of the coefficients described for nutrients. The coefficient

HYPE Model Documentation Online HYPE documentation land-use dependent parameter *ocsoilslp*, and the coefficient *satact* is replaced by land-use dependent parameter *ocsoilsat*.



Production of humusC from fastC

Some of the litter fall (added to fastC) is converted into humus. For HYPE this means that fastC (where litter fall was added) is transformed to humusC in the uppermost soil layer. In Figure 2 the pool of humusC is denoted slowC because of its slower transformation rates.

The other soil layers (k) also have a transition from fastC to humusC. The loss of fastC does not all end up in the humusC pool, but a proportion (parameter minc) is mineralized in the process. The transformation (fasttohumus, mg/m2/d) depends on soil moisture (smfcn) and temperature (tmpfcn), amount of Oc in the pool (fastC) and a vegetation dependent parameter klh.

$$fast to humus\left(k\right) = klh \times tmp \, fcn\left(k\right) \times sm \, fcn\left(k\right) \times fast C\left(k\right)$$

Turnover of fastC

Turnover of fastC is a sink for fastC and a source of dissolved OC in soil water in all soil layers (k = 1-3). The loss of fastC does not all go to the OC, but a proportion (parameter minc) is mineralized. Turnover (transfC, mg/m2/d) depends on a general parameter (klo), the temperature function (tempfcn), humidity function (smfcn) and the pool of fastC (fastC).

$$transfC(k)=klo \times tempfcn(k) \times smfcn(k) \times fastC(k)$$

In dry conditions a flow in the opposite direction can also occur. The transformation of OC to fastC is a decrease of OC and a source of fastC in all soil layers (k = 1-3). The loss of OC is not all to fastC but a proportion (parameter minc) is mineralized. Turnover (doctofast, mg/m2/d) depends on a general parameter (kof) and the pool of OC (OCpool). The flow is limited that the soil layer temperature must be less than 5 °C, the soil moisture (sm) must be less than field capacity and moisture function (smfcn) must be less than the parameter koflim.

$$doctofast(k) = kof \times OCpool(k)$$

Turnover of humusC

Turnover of humusC is a sink for humusC and a source of OC in all soil layers (k = 1.3). The loss of humusC does not all go to the DOC, but a proportion (parameter *minc*) is mineralized. Turnover (*transhC*, mg/m2/d) depends on a general parameter (*kho*), temperature function (*tempfcn*), humidity function (*smfcn*) and the pool of humusC (*humusC*).

$$transhC\left(k\right)\!\!=\!kho\!\times\!tempfcn\left(k\right)\!\!\times\!smfcn\left(k\right)\!\!\times\!humusC\left(k\right)$$

Percolation

Organic carbon is lost from the soil water as it flows down through the soil layers and where it is dissipated to become a regional groundwater flow. The decrease in concentration depends on soil moisture and temperature and a calibration parameter.

$$conc \!=\! conc \!\times \! \Big(1 \text{-} par \!\times\! tmp \, fcn \!\times\! sm \, fcn \,\, \Big)$$

The soil moisture function and temperature function are the general functions described for soil processes. Percolation uses the coefficients for soil moisture function, not the parameters as the transformations. The parameter, *par* in the equation, is called *kcgwreg* for regional groundwater flow formation and *koc* for percolation between soil layers. Both are general parameters.

Links to relevant procedures in the code

Modules (file)	Procedures	Section
npc_soil_processes (npc_soil_proc.f90)	initiate_soil_npc	initial values
	soil_carbon_processes	production of humusC from
	soil_carbon_pool_transformations	fastC, turnover
	doc_percolation_reduction	percolation

Riparian zone

Runoff from soil layers may flows through a riparian zone before it reaches the local river. Surface runoff and drainage water from drainage pipes reaches the local river without passing through the riparian zone. In the riparian zone the levels of OC are affected, while flows remain unchanged. The change depends on soil temperature, class altitude (elev (in masl)), the water table (gwat) and its recent change, season and soil moisture (sm). The runoff concentration (conc(i)) of each soillayer (k) increases with the factor:

$$f\!\left(k\right)\!\!=\!1\!+ripz\!\times\!tmpfcn\left(k\right)\!\!\times\!\!\left(\frac{elev}{100}\right)\!\!\times\!f_{g\!r\!w}\!\times\!f_{season}\!\times\!f_{sm}$$

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$$conc(k)=f \times conc(k), k=1..3$$

The temperature function (*tmpfcn*) is the usual of soil processes (see above). The following equations describe the other process functions:

$$\begin{split} \boldsymbol{f}_{grw} &= e^{ripe \times gwat} \\ \boldsymbol{f}_{season} &= \begin{cases} rips & autumn \\ 1 & otherwise \end{cases} \end{split}$$

$$\boldsymbol{f}_{sm} \! = \! \left\{ \begin{array}{ll} 0 & \! sm \! \leqslant \! wp \\ \boldsymbol{f}_2 \! \left(sm \right) & \! wp \! < \! sm \! < \! pw \\ satact & \! sm \! \geqslant \! pw \end{array} \right. \label{eq:fsm}$$

$$f_{2}\!\left(sm\right)\!\!=\!\!\left\{\begin{array}{l} \min\!\left(1,satact\!+\!\frac{\left(1\!-\!satact\right)\!\!\times\!pw\!-\!sm}{d\!\times\!\Theta_{upp}},\!\frac{sm\!-\!wp}{d\!\times\!\Theta_{low}}\right) & rising\,grw\\ \min\!\left(1,satact\!+\!\frac{\left(1\!-\!satact\right)\!\!\times\!pw\!-\!sm}{d\!\times\!\Theta_{upp}},\!\frac{satact\!\times\!sm\!-\!wp}{d\!\times\!\Theta_{low}}\right) & sinking\,grw \end{array}\right.$$

The activation of riparian zone processes is based on land use. The land use dependent parameter *ripz* determines the overall level of increase in concentration in the riparian zone, and if set to zero no riparian zone processes are used. In addition two general parameters can influence the effect of the riparian processes; *ripe* which determines the groundwater level dependence, and *rips* which determines the seasonal influence. Season division is determined by ten-day and twenty-day averages of air temperature (T10, T20). Autumn is defined as the period when T10 is less than T20. The soil moisture function is different for an increasing (rising) and sinking ground water table (figure 2). It contains coefficients $\Theta_{upp} = 0.12$ $\Theta_{low} = 0.08$ and saturation (*satact* = 0.6). It depends on the soil moisture of all layers together (*sm*) and the water-holding capacity parameters; *wp* - wilting border and *pw* - total pore volume, in fractions of total soil layer thickness (*d*).

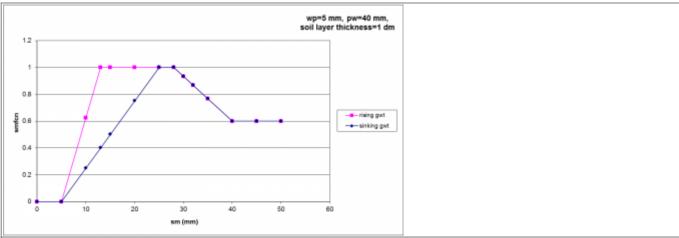


Figure 2: Example of riparian zone soil moisture function, and the dependence on changes in the groundwater levels.

Links to relevant procedures in the code

Modules (file)	Procedures	
npc_soil_processes (npc_soil_proc.f90)	class_riparian_zone_processes	
	riparian_moisture_factor	

Rivers and lakes

Primary production and mineralization

Primary production is a source of organic carbon in rivers and lakes, while mineralization is a sink. Primary production and mineralization is calculated the same way as for nitrogen, but with its own calibration parameter (*wprodc*). The potential carbon transformation (*minprodCpot*, kg / day) is proportional to the potential nitrogen transformation (*minprodNpot*, see NP section) with a transformation rate that depends on the carbon-nitrogen ratio (*NCratio* = 5.7). The calculated mineralization of organic carbon is limited to a maximum of 50% of the available OC pool. If phosphorus is not modelled a long-term average total phosphorus concentration as a lake region dependent parameter (*tpmean*) is used. If set, the long-term average concentration is reduced by the general parameter *limsedPP* before using it in the concentration function.

$$tmpfcn_1 = \frac{watertemp}{20}$$
.

$$tmpfcn_2 = \frac{\left(T_{10} - T_{20}\right)}{5}$$

 $tmpfcn = tmpfcn1 \times tmpfcn2$

$$TPfcn = \frac{\left(TPconc \cdot limsedPP\right)}{\left(TPconc \cdot limsedPP + halfsatTPwater\right)}$$

 $minprodNpot = wprodc \times TPfcn \times tmpfcn \times area$

 $minprodCpot = minprodNpot \times NCratio$

Sedimentation

Sedimentation in lakes is a sink for OC and works the same way as for organic nitrogen and particulate phosphorus. Sedimentation (*sedOC*, *kg/day*) is calculated as a function of water concentration and lake area (*area*). The parameter *sedoc* is general or can be specified for each lake.

 $sedOC = sedoc \times waterconcOC \times area$

Links to relevant procedures in the code

Modules (file)	Procedures	Sections	
	oc_processes_in_river	primary production and mineralization	
npc_surfacewater_processes (npc_sw_proc.f90)	oc_production_mineralisation		
	calculate_lake_tpmean		
	calculate_river_tpmean		
	oc_processes_in_lake	primary production and mineralization	
		sedimentation	
	oc_sedimentation	sedimentation	

Links to file reference

Section	Symbol	Parameter/Data	File	
Sources of		resc	CropData.txt	
organic material		litterdays	par.txt	
Soil processes		humusc1, humusc2, humusc3, fastc1, fastc2, fastc3		
	θ_{low}	ocsoilslp or 0.08		
	satact	ocsoilsat or 0.6		
	minc, klh, klo, kof, kho	minc, klh, klo, kof, kho		
		koflim		
	par	kcgwreg or koc		
Riparian zone	elev	calculated from mean_elev and dhslc_nn	GeoData.txt	
	ripz, ripe, rips	ripz, ripe, rips		
	wp	calculated from wcwp, wcwp1, wcwp2, wcwp3	par.txt	
	pw	calculated from wcwp, wcwp1-wcwp3, wcfc, wcfc1-wcfc3, wcep, wcep1-wcep3	μαι.ιχι	
	d		GeoClass.txt	
Rivers and lakes	area		GeoData.txt	
	wprodc, limsedpp, sedoc	wprodc, limsedpp, sedoc	par.txt	
		tpmean		
	halfsatTPwater	hsatTP]	

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